VII. BOTTOM SEDIMENTS IN THE GH78-1 AREA

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Method

31 sediment samples were obtained from the survey area with Okean-70 grab sampler or box sampler, and 1 sample with deep sea camera in its frame pipe. Those samples were treated in the same process as those in the previous cruise, GH77-1 (NAKAO, 1979).

Composition of the coarse fraction and the types of sediment

Natural property and composition of the coarse fraction (>63 micron) settled in the sedimentation tube were described in the previous report (NAKAO, 1979).

The following types of sediments, two of which, i.e. (6) and (7) are newly defined in this report, are identified;

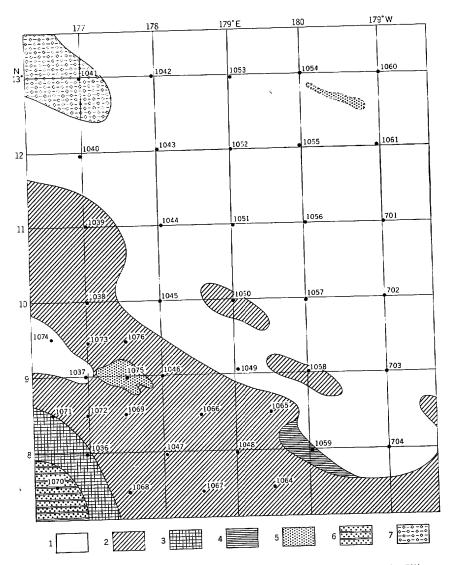
- (1) deep sea clay: coarse fraction <10%, or 10-30% and non-biogenic (consists largely of "mineral" and silicified radiolarian remains),
- (2) siliceous clay: coarse fraction 10-30%, radiolarian,
- (3) calcareous-siliceous clay: coarse fraction 10-30%, foraminiferal-radiolarian,
- (4) siliceous ooze: coarse fraction >30%, radiolarian,
- (5) calcareous ooze: coarse fraction >30%, foraminiferal,
- (6) calcareous-siliceous ooze: coarse fraction >30%, foraminiferal-radiolarian, and
- (7) zeolitic mud: coarse fraction >30%, zeolitic.

Distribution of surface sediment

Fig, VII-1 shows distribution of the surface sediments classified into seven types mentioned in the preceding section. The southern half of the area is characterized by the distribution of biogenic sadiments, on the other hand the northern half by non-biogenic ones.

The coarse fraction consists almost exclusively of foraminiferal remains at the station 1075 (water depth, 4,028 m), and foraminiferal-radiolarian at St. 1070 (WD 4,947 m) as shown in Fig. VII-2(h). Maximum water depth where foraminiferal remains are found in the coarse fraction of the surface sediment is 5,250 m (St. 1067).

The main zone of siliceous clay (St. 1059) is continuous from the east adjacent area, GH77-1. Siliceous ooze, the coarse fraction of which consists dominantly of stained radiolarian remains, is distributed in the borderland between siliceous clay and deep sea clay, in the same manner as the case in GH77-1 (St. 731 and 732). In other words, the change of the coarse fraction ratio is not concordant with the change of water depth. This fact suggests some sorting effect which can concentrate radiolarian remains in the sediment.



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Fig. VII-1 Distribution of the surface sediments. 1. Deep sea clay, 2. Siliceous clay, 3. Calcareous-siliceous clay, 4. Siliceous ooze, 5. Calcareous ooze, 6. Calcareous-siliceous ooze, 7. Zeolitic mud.

Deep sea clay is extensively distributed. However, it is very important to point out that more than 30% of the coarse fraction of deep sea clay consists of radiolarian (including the stained) remains, at about half the number of those stations (Fig. VII-3 (a) and (b)). At the remaining stations, the coarse fraction of deep sea clay consists of dominant zeolite (phillipsite) grains and abundant or common silicified radiolarian remains, in general.

There is presumable distribution of calcareous ooze near the northeast corner

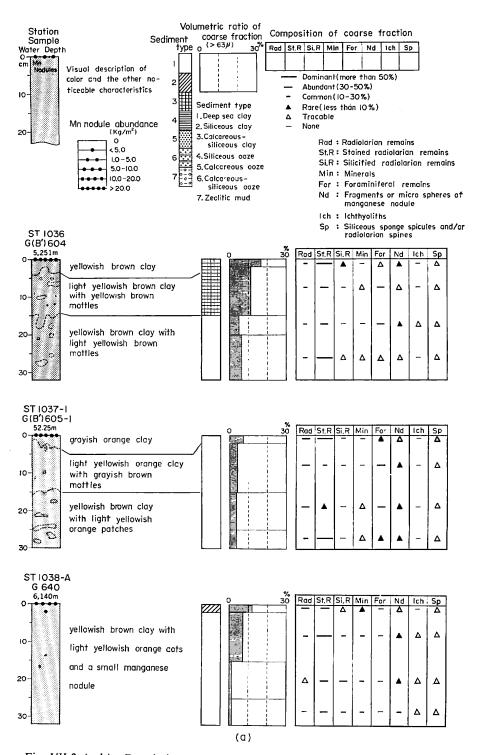
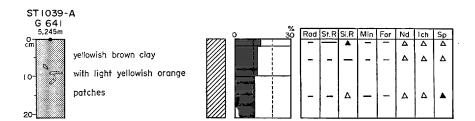
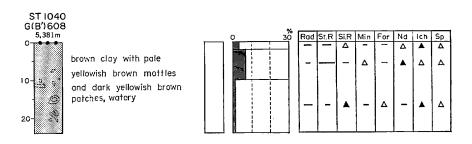
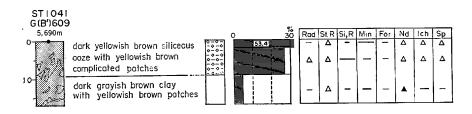
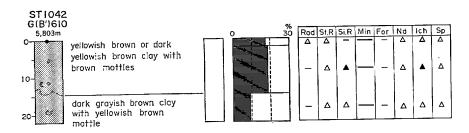


Fig. VII-2 (a-h) Description of the samples from Okean-70 or box sampler.

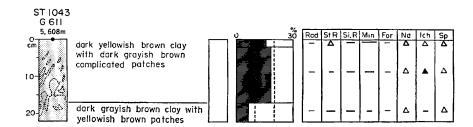


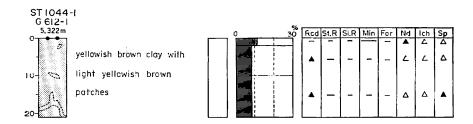


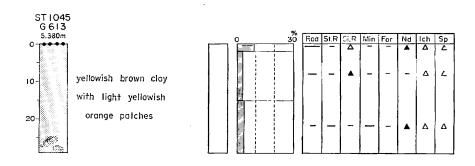


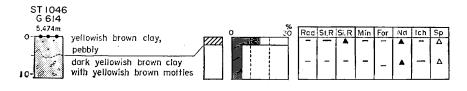


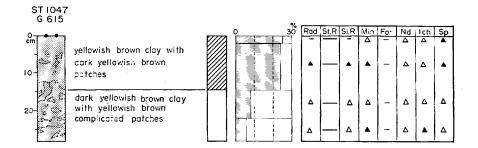
(b)

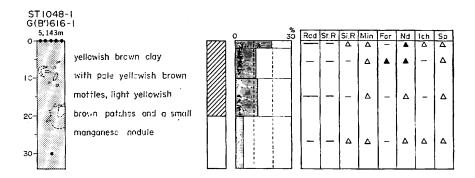


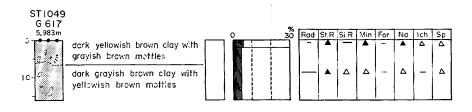


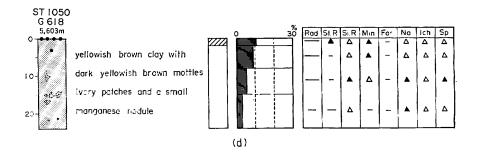


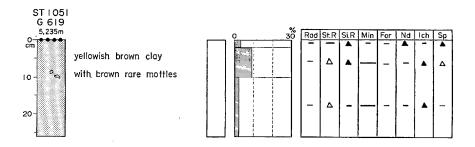


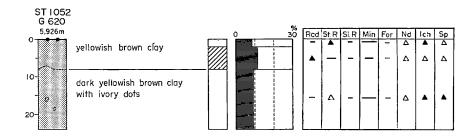


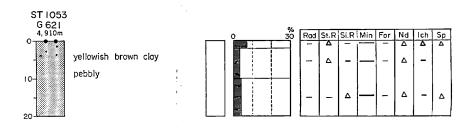


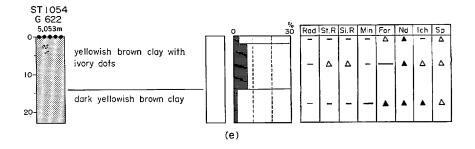


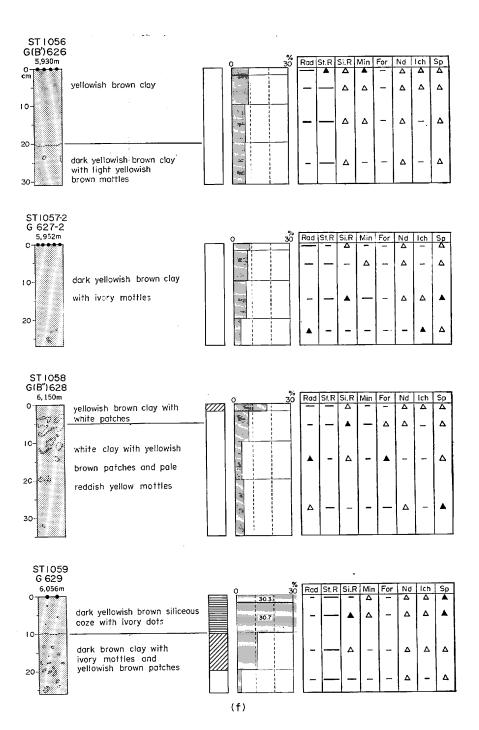


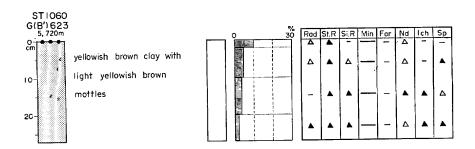


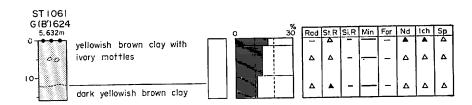


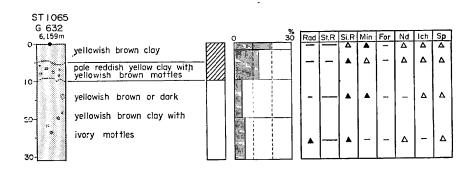


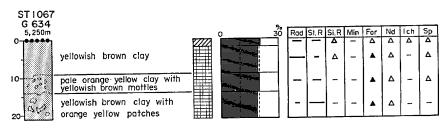




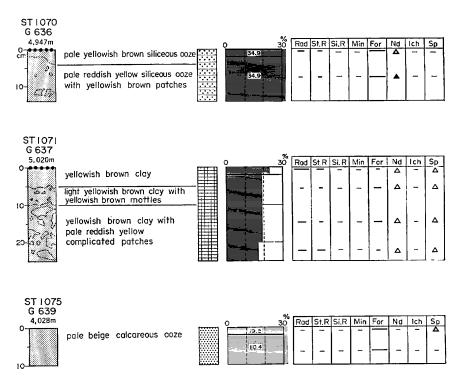




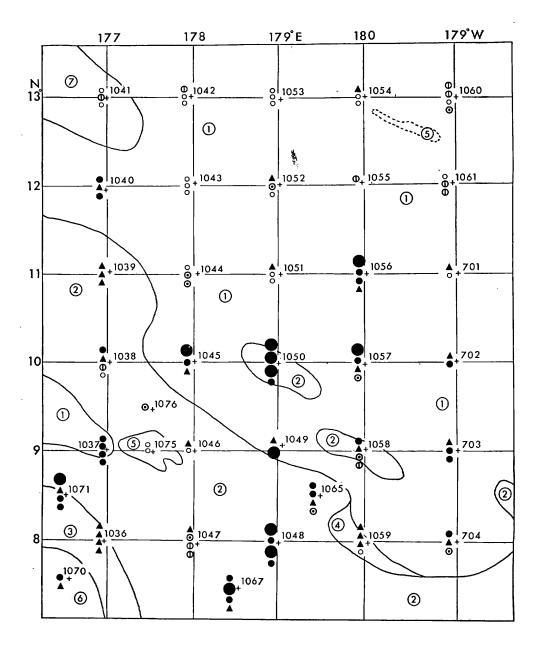




(G)

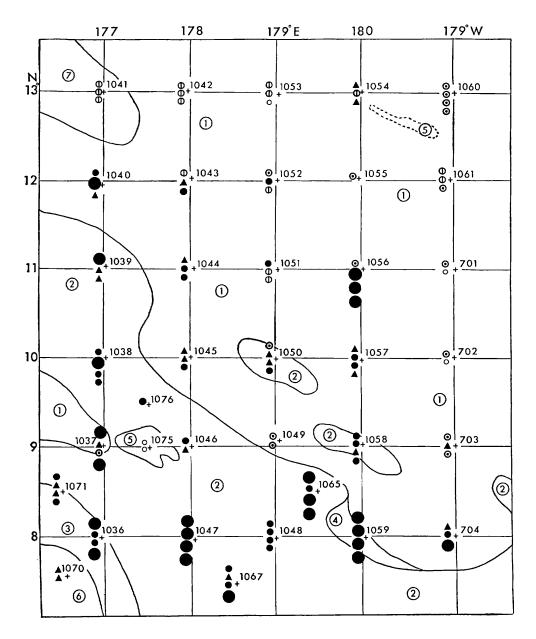


(h)

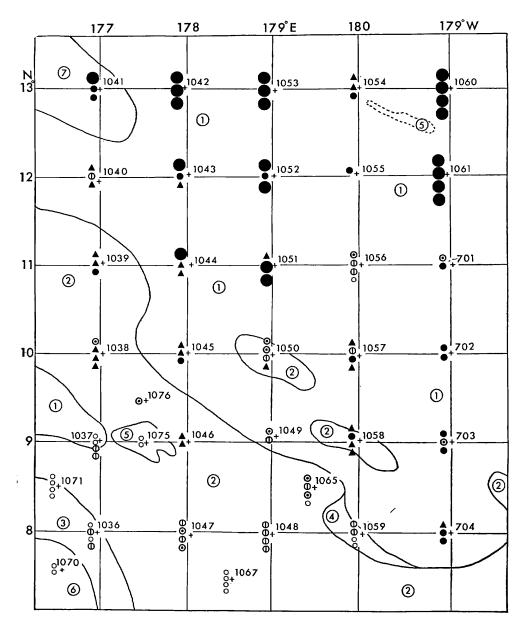


Dominant ●Abundant ▲ Common ⊕ Rare ⊕ Tracable ◦ None (a)

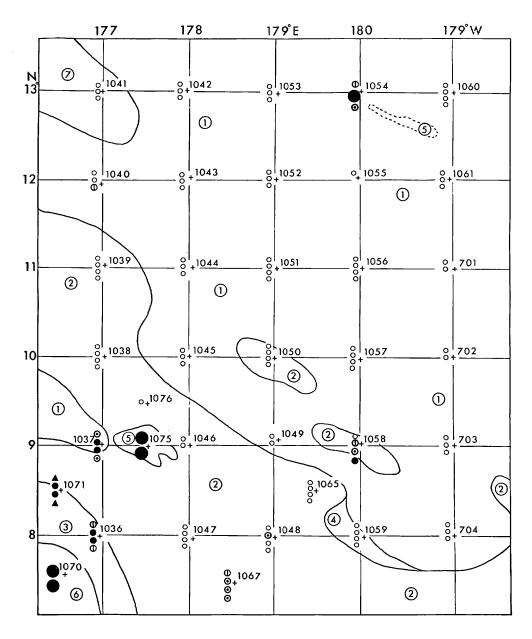
Fig. VII-3 (a-d) Distribution of the major components in the coarse fraction (>63 μm). (a) Radiolarian remains, (b) Stained radiolarian remains, (c) Minerals and (d) Foraminiferal remains. Expression of abundance corresponds to that in Fig. VII-2, namely "Dominant" means the abundance of more than 50%, "Abundant" between 30 and 50% etc.



Dominant ●Abundant ▲Common ©Rare ⊕Tracable oNone (b)



Dominant ◆Abundant ▲ Common ⊗ Rare ⊕ Tracable ∘ None
(c)



■ Dominant ◆Abundant ▲ Common ⊕ Rare Φ Tracable ◦ None
(d)

of the survey area. This presumption is based upon dominant foraminiferal remains in the coarse fraction of the sample just beneath the surface from St. 1054 (Fig. VII-3 (d)) and topographic high shallower than 4,500 m.

Zeolitic mud from St. 1041 has the coarse fraction which consists of dominant (about 60%) zeolite (phillipsite) grains and common (almost 30%) silicified radiolarian remains. The coarse fraction ratio of that sample is up to 53.4%.

Sequence in sediment cores

Fig. VII-2 (a-h) shows vertical section, type of sediment, composition of the coarse fraction of the cores from Okean-70 or box sampler. Surface and vertical distribution of major components in the coarse fraction is shown in Fig. VII-3 (a-d), photographs of vertical section of the core samples in Fig. VII-4 (a-e) and Softex-radiographs of sliced core samples in Fig. VII-5 (a-f).

Mottle structure, which suggests bioturbation in the past, is more or less observed on the vertical sections of almost all cores. However, occurrence of abundant and large-scale mottle structure is restricted in the zone south of the parallel of latitude, 9°N.

The sediment with the coarse fraction of more than 10% is thin in general. Those have such fraction ratio throughout the whole core length are found at the stations, 1039 (core length: 20 cm), 1061 (16 cm), 1067 (21 cm), 1070 (14 cm), 1071 (25 cm) and 1075 (10 cm). Among those, however, only the two samples, i.e. those from St. 1039 and 1061 are free from foraminiferal remains in the coarse fraction. The sample from St. 1039 has the coarse fraction consisting of dominant (ca. 60%) stained radiolarian remains in the surface and abundant (ca. 45%) "mineral" in the bottom. And the coarse fraction of the sample from St. 1061 consists dominantly of zeolite grains.

Abrupt decreasing (to less than half of that in the uppermost or upper adjacent part) of the coarse fraction ratio is observed in about half the number of core samples. This phenomenon is common also in some samples, the coarse fractions of which consist largely of "mineral".

In some vertical sequences, temporary increasing of foraminiferal remains in the coarse fraction is observed (Fig. VII-3(d), St. 1036, 1037, 1048, 1054 and 1071). It cannot be clarified whether this increasing is caused by the oscilation of primary supply of calcareous material or rather rapid redistribution of pre-existent foraminiferal sediment.

Stratigraphy

Although concrete stratigraphic evidence has not been proved for this cruise, some presumption may be given through the petrographic comparison of the samples with those in the previous cruise, GH77-1, for which radiolarian biostratigraphy has been clarified for 8 samples including 3 long cores (SAKAI, private communication).

Compiled data of that radiolarian biostratigraphy, foraminiferal biostratigraphy (NATORI, private communication, in NAKAO, 1979) and the coarse fraction petrography suggest following criteria for presumption of geologic age of the samples.

- The sample, coarse fraction of which consists fresh radiolarian remains as a major component, small amount of stained radiolarian remains, if any, and no silicified radiolarian remains, will be of Quaternary age. Volumetric ratio of the coarse fraction of such a sample usually exceeds several percent.
- 2) Calcareous and calcareous-siliceous sediments which are distributed at topographic highs in the area of GH77-1 are of Quaternary age.
- 3) Abrupt decreasing of the coarse fraction ratio mentioned above, in the sample with radiolarian and stained radiolarian remains in its coarse fraction, seems to appear around the boundary between Tertiary and Quaternary deposits.
- 4) Silicified radiolarian remains indicate older age, in general, but reworking and redistribution of those particles, and bioturbation of the sediments should be taken into account.
- 5) "Mineral" grains (mainly zeolite mineral) increase in the coarse fraction toward older age. And the sediment, in the coarse fraction of which "mineral" component is abundant or dominant, seems to be of Tertiary age. Considerations for reworking and redistribution of those, and bioturbation of the sediments are necessary, too.

According to the application of those criteria mentioned above, following presumptions on geologic age and stratigraphy of the sediment in this area (GH78-1) can be given.

- 1) The surface sediments, which are distributed in the zone south of the parallel of latitude 10°N, will be of Quaternary age, except at St. 1059 and 1076, where the sediments may be of Tertiary age. However, the thickness of the Quaternary sediment seems to be thinner than 20 cm in the case of non-calcareous sediment, because the sediment showing characteristics of Tertiary age underlies those Quaternary sediment in such a case.
- 2) In the northern half of the area, the sediments of possible Quaternary age are distributed at the stations 1051, 1052, 1054 and 1056. Other than those it seems Quaternary sediments are distributed on the surface of the stations 1039, 1040 and 1055.

As for the zeolite rich deep sea clay and the zeolitic mud in the northern half of the area, further study is necessary to presume the geologic age of those. Kastner and Stonecipher (1978) point out that phillipsite predominates in volcanogenic and clayey sediments (sedimentation rates of $< 5 \text{ m}/10^6$ years) of Recent to Eocene age, that it starts to form at the sediment-water interface and that the most common precursor is volcanic (basaltic) material, in particular basaltic glass and palagonite.

Thus, the occurrence of phillipsite is not to be a criterion to distinguish the geologic age, Quaternary from Tertiary, of the sediment. However, phillipsite occurs with silicified radiolarian remains in this area, in general. Furthermore, phillipsite rich sediment underlies calcareous sediment which will be of Quaternary age at the station, 1054. Those facts lead us to the presumption that the phillipsite rich sediment in this area is not to be of Quaternary age.

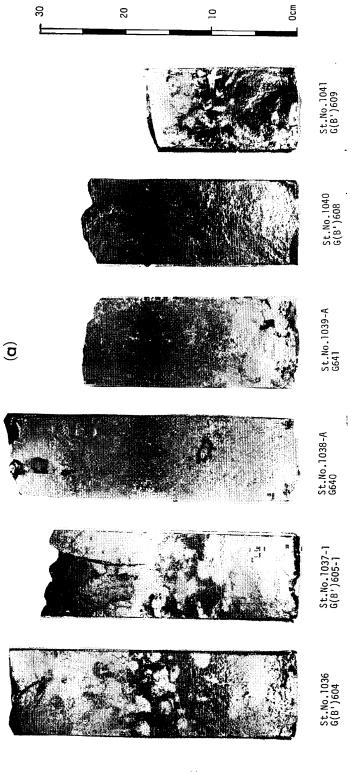
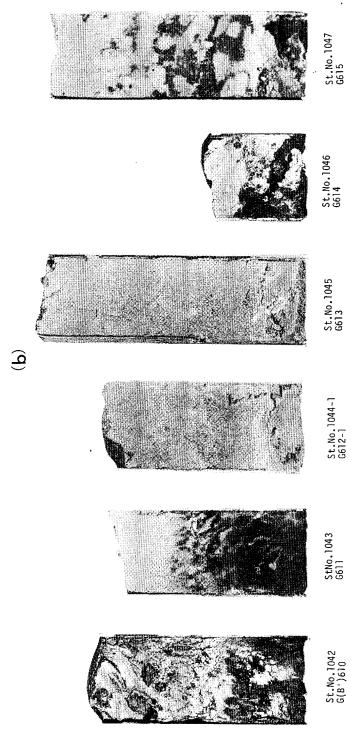
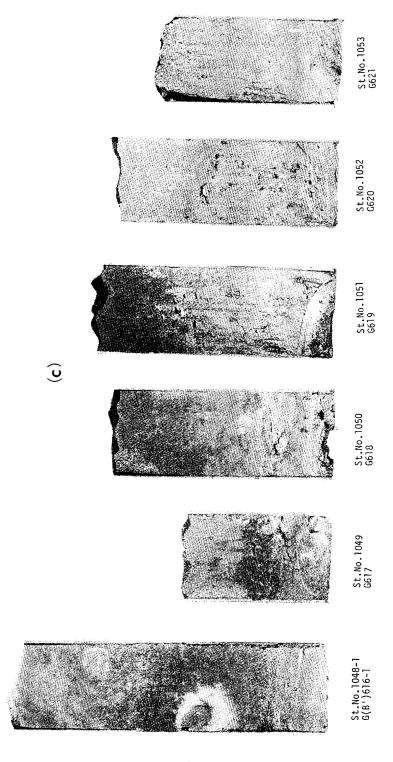
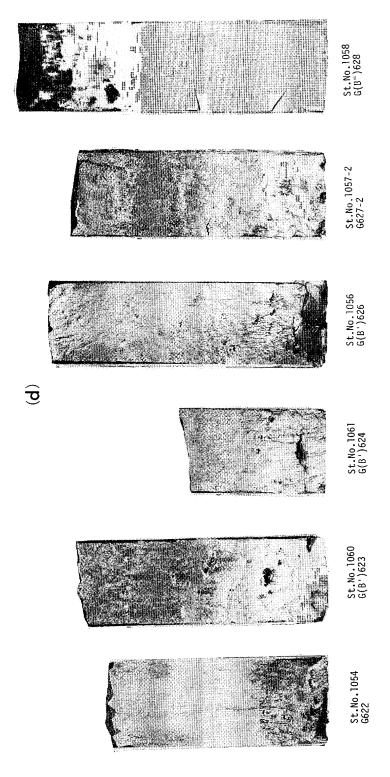
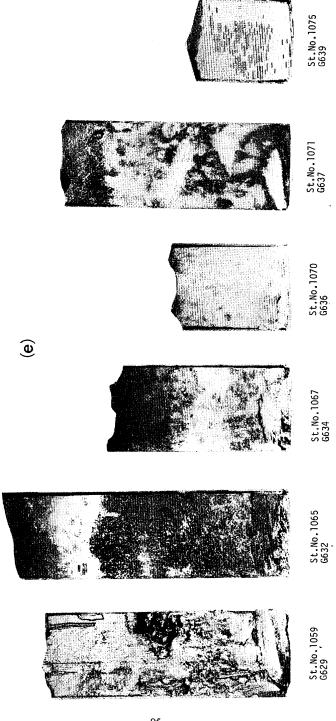


Fig. VII-4 (a-e) Photographs of vertical section of the core samples from Okean-70 or box sampler.









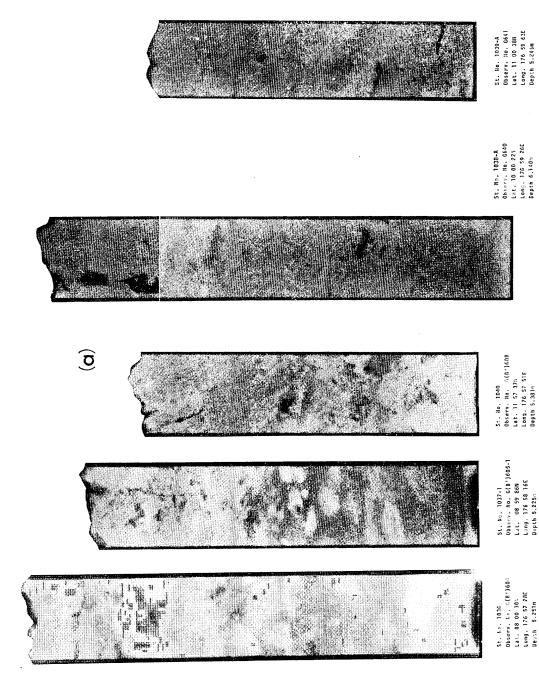
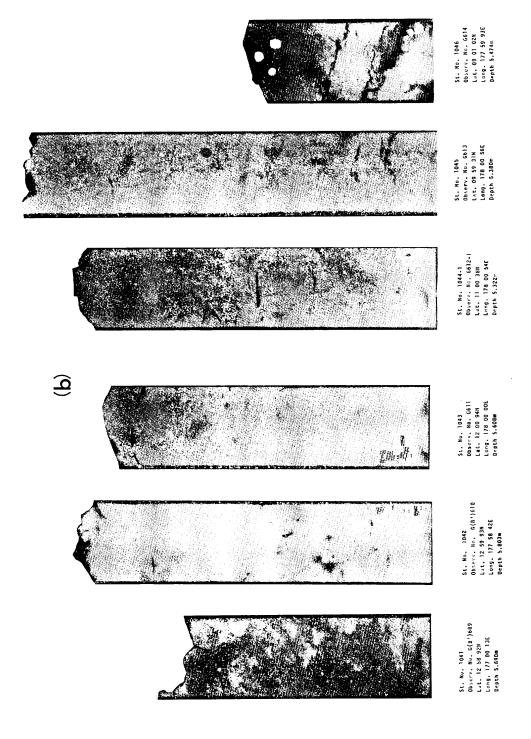
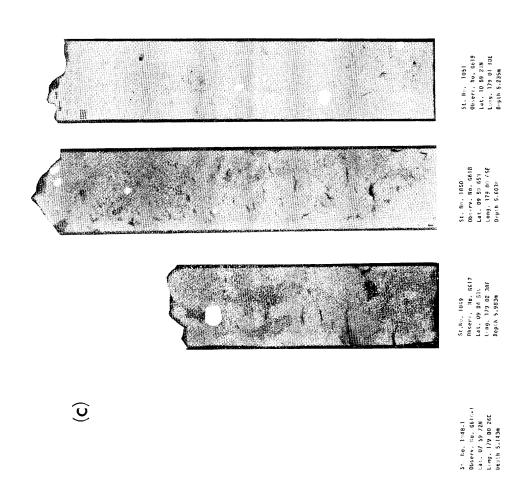


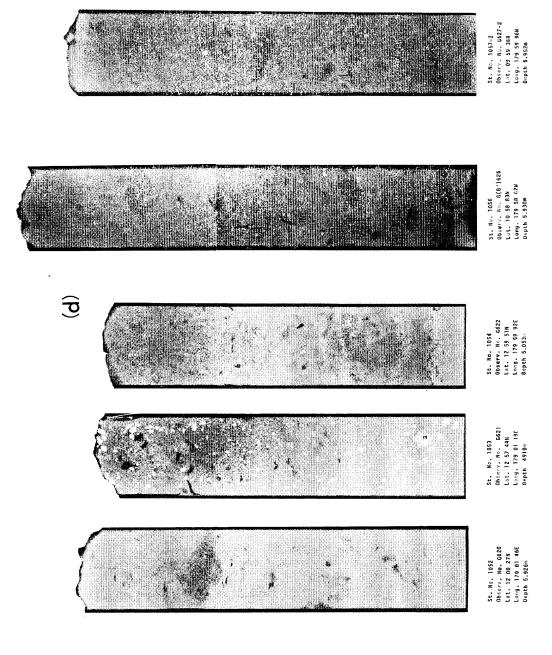
Fig. VII-5 (a-f) Softex-radiographs of sliced core samples from Okean-70 or box sampler.

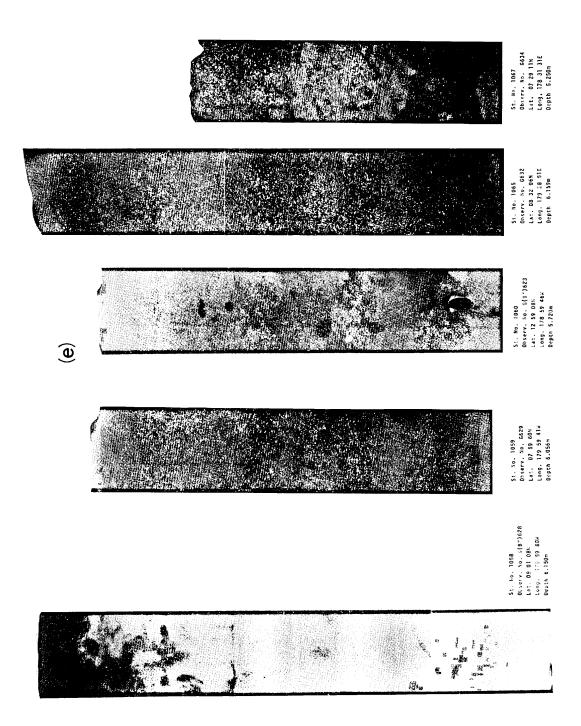






St.No. 1047 Observ. ho. G615 Lat. 07 59 15N Long. 178 01 67f 'epth 5.660m





2 Observ. No. G(8')624 Ldt. 11 59 82H Long. 179 02 60W Depth 5.632m St. No. 1061 $\widehat{\mathbf{+}}$ St. No. 1075 Observ, No. 6639 Lat. 09 00 17N Long. 177 30 09E Depth 4.02 Um 5t, No. 1071 Observ, No. 6637 Lat. 08 30 50H Long. 176 30 39E Umpth 5.020m St. Ho. 1070 Observ. No. Godb Lat. O7 33 83N Long. 176 32 97E Uspth 4,947m

Interrelation between distribution of manganese nodules and sediments

- 1) Manganese nodule is none or scarse on calcareous or siliceous ooze.
- 2) Manganese nodules occur abundantly ($> 10 \text{ kg/m}^2$) only with sediment of possible Quaternary age.
- 3) The highest abundance ($> 20 \text{ kg/m}^2$) of manganese nodules occur with the sediment, the coarse fraction of which is largely radiolarian or foraminiferal-radiolarian, except at the station 1054.

References

- KASTNER, M. and STONECIPHER, S. A. (1978) Zeolites in pelagic sediments of the Atlantic, Pacific, and Indian Oceans, *In* L. B. SAND and F. A. MUMPTON (eds.), *Natural Zeolites*, Pergamon Press, p. 199–220.
- NAKAO, S. (1979) Bottom sediments, In T. MORITANI (ed.), Geol. Surv. Japan Cruise Rept., no. 12, p. 131–151.