Report

Late Triassic radiolarians and conodonts from a chert pebble within the Lower Pleistocene Higashihigasa Formation of the Kazusa Group, Boso Peninsula, Japan

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ITO Tsuyoshi, MUTO Shun and UTSUNOMIYA Masayuki (2022) Late Triassic radiolarians and conodonts from a chert pebble within the Lower Pleistocene Higashihigasa Formation of the Kazusa Group, Boso Peninsula, Japan. *Bulletin of the Geological Survey of Japan*, vol. 73(3), p. 93–101, 6 figs and 1 table.

Abstract: Radiolarians and conodonts were obtained from a chert pebble within the conglomerate of the Higashihigasa Formation, Kazusa Group, Boso Peninsula, Japan. Based on the occurrence of radiolarians (*Praemesosaturnalis* sp. cf. *P. heilongjiangensis*) and conodonts (*Mockina* sp.), the chert pebble is considered to be Late Triassic (middle to late Norian) in age. This chert pebble is presumably derived from a Jurassic accretionary complex distributed in its provenance.

Keywords: radiolarian, conodont, Triassic, Pleistocene, Kazusa Group, conglomerate, provenance, Jurassic accretionary complex, Boso Peninsula, Chiba Prefecture

1. Introduction

Clasts within sediments, such as conglomerate, are supplied from the surrounding geologic units and record information of the provenance. Microfossils including radiolarians and conodonts can assign the age of the clasts, even if they are small clasts.

Radiolarian-bearing clasts have been reported from the Paleozoic to Cenozoic (Table 1) and compiled (e.g. Ishida *et al.*, 2003; Ito *et al.*, 2017a, e), but radiolarian-bearing clasts within the Quaternary are poorly investigated: only by Ito *et al.* (2020) as far as we know. Ito *et al.* (2020) indicated the presence of a water system different from the present one in the Nishi-Mikawa region, central Japan, based on microfossil-bearing clasts within the Quaternary in several areas and horizons will contribute to reconstruction of changes of provenances and water systems.

The Pleistocene Kazusa Group is distributed in the Boso Peninsula (Fig. 1). Some formations of the group include conglomerate layers. We investigated the conglomerate for the accumulation of data of microfossil-bearing clasts within the Quaternary. We consequently discovered radiolarians and conodonts from a chert pebble within the conglomerate of the Higashihigasa Formation of the Pleistocene Kazusa Group, Boso Peninsula (Fig. 1). In this article, we note the microfossils as the first report of microfossil-bearing clasts within the Quaternary in the Boso Peninsula.

2. Geologic setting

The Kazusa Group is mainly composed of shallow- to deep-marine successions (>3,000 m in total thickness). The group generally comprises the Kurotaki, Katsuura, Namihana, Ohara, Tomiya, Kiwada, Otadai, Umegase, Higashihigasa, Kokumoto, Kakinokidai, Ichijiku, Chonan, Mandano, Kasamori and Kongochi formations (Tokuhashi and Endo, 1984; Nakajima and Watanabe, 2005; Utsunomiya and Ooi, 2019) (Fig. 2).

The microfossil-bearing pebble dealt with in this study was collected from the conglomerate of the Higashihigasa Formation. The Higashihigasa Formation is considered as a canyon-fill deposit which interfingers with the submarine-fan deposits defined as the Otadai and Umegase formations (Yamauchi *et al.*, 1990). Chronostratigraphy based on magneto-, tephro- and bio-stratigraphy suggests the Otadai and Umegase formations were deposited during the Early Pleistocene (Calabrian) (Kazaoka *et al.*, 2015). The clast-bearing conglomerate is intercalated above the U10 tephra bed, which is near the Marine Isotope Stage 24 (Pickering *et al.*, 1999), at about 0.9 Ma.

The sample locality is along a tributary of the Obitsu River, north of Mt. Otsuka (Fig. 3). It is located in Otomi, Kimitsu City, Chiba Prefecture in administrative division. The chert clasts occur at the basal part of a conglomerate (2 m in thickness), associated with bioclasts (e.g. mollusks) and other kinds of gravels such as sedimentary rocks and volcanic rocks (Fig. 4A).

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| Erathem | System | Series | Reference |
|------------|---------------|----------------------|--------------------------------------|
| Cenozoic | Quaternary | Pleistocene | Ito et al. (2020) |
| | | | Ito et al. (2022) [this study] |
| | Neogene | Pliocene? | Matsuoka (1998) |
| | | | Yamamoto et al. (2012) |
| | | Pliocene | Kawajiri and Kashiwagi (2012) |
| | | | Kashiwagi et al. (2013) |
| | | | Utagawa et al. (2017) |
| | | Miocene | Umeda et al. (1992) |
| | | | Kashiwagi (2012) |
| | | | Ito and Nakamura (2021) |
| | | | Yabuta et al. (2021) |
| | Paleogene | Oligocene | Umeda (1997) |
| | | Eocene? | Kamemura and Okamura (1994) |
| | | Paleocene? | Kishu Shimanto Research Group (2017) |
| Mesozoic | Cretaceous | Upper | Suzuki et al. (1996) |
| | | | Inose et al. (2018) |
| | | Lower | Kojima (1986) |
| | | | Takeuchi et al. (1991) |
| | | | Umeda et al. (1995) |
| | | | Ishida and Hashimoto (1997) |
| | | | Ishida (1999) |
| | | | Umeda and Sugiyama (1998) |
| | | | Matsukawa and Takahashi (1999) |
| | | | Nikaido and Matsuoka (2004) |
| | | | Tomita et al. (2007) |
| | | | Ito et al. (2012) |
| | | | Ito et al. (2014) |
| | | | Ito et al. (2015) |
| | | | Kashiwagi and Isaji (2015) |
| | | | Takeuchi et al. (2015) |
| | | | Ozeki et al. (2021) |
| | Jurassic | Upper | Saida (1987) |
| | | Middle? | Ito et al. (2016) |
| | | | Ito et al. (2017b) |
| | | Lower | Kumazaki and Kojima (1996) |
| | | | Ito <i>et al.</i> (in press) |
| | Triassic | Upper | Kametaka (1997) |
| | Triassic? | ** | Saito and Tsukamoto (1993) |
| | | | Kamata (1997) |
| Paleozoic | Permian | Gudalupian–Lopingian | Matsuoka and Kuwahara (2021) |
| | | 1 1 8 | Takemura <i>et al.</i> (1996) |
| | Carboniferous | Mississippian | Uchino and Kurihara (2019) |
| Paleozoic? | | 11 | Ito <i>et al.</i> (2017c) |
| | | | Ito <i>et al.</i> (2017d) |

Table 1 Major previous studies of radiolarian-bearing clasts.

3. Method

Four pebbles collected from the conglomerate were processed with the following method to extract microfossils. They were crushed into some fragments to create more surface area. The crushed pebbles were soaked in 5 % hydrofluoric acid at room temperature, about 20–25 °C, for 24 h. The residues were collected by a sieve with

a mesh opening of 0.054 mm. This process was repeated four times. Fossil specimens in the residues were picked up and mounted on stabs. The specimens on the stabs were photographed by scanning electron microscopy. Part of the residues was enclosed within a slide prepared with a photocrosslinkable mounting medium (GJ-4006, Gluelabo Ltd.). The slides were photographed using a transmitted light microscope.

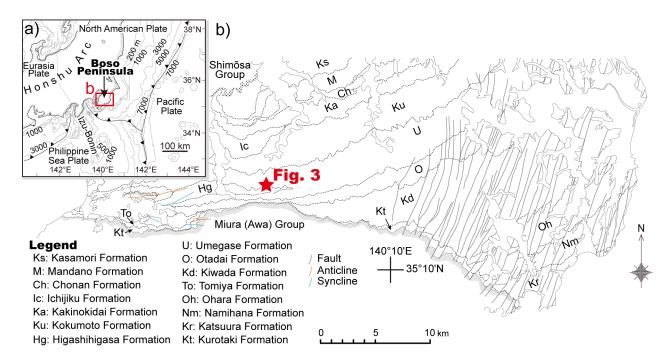


Fig. 1 Index and simplified geologic maps of Boso Peninsula (modified after from Utsunomiya et al., 2019).

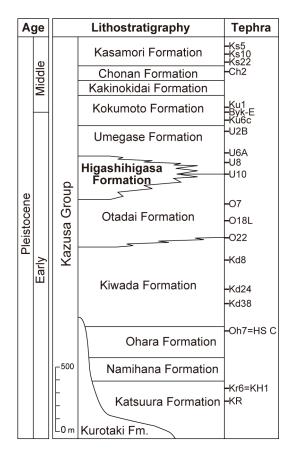


Fig. 2 Stratigraphy of the Kazusa Group in the Boso Peninsula (modified after from Utsunomiya *et al.*, 2019).

4. Microfossil occurrences

Among the four pebbles, only one pebble (sample 21112309a) yielded radiolarians and conodonts. The pebble is a rounded reddish chert, and its diameter is about 3 cm (Fig. 4B). The radiolarian and conodont specimens are shown in Fig. 5.

Some radiolarians, *Praemesosaturnalis* sp. cf. *P. heilongjiangensis* Yang and Mizutani, *Praemesosaturnalis*? sp., *Paroertlispongus*? sp., Poulpidae? gen. et sp. indet. and Spumellaria gen. et sp. indet. were extracted. The specimens identified as *Praemesosaturnalis* sp. cf. *P. heilongjiangensis* (Figs. 5.1–5.5) seem to have spines with an elevated margin and bifurcated end. Such a spine is known in a broken specimen of *Praemesosaturnalis heilongjiangensis* (Yang and Mizutani, 1991).

Two specimens of conodonts were extracted. One specimen (Fig. 5.33) can be identified as *Mockina* sp. due to the following characters: a carina extending to the posterior end, a platform with only one sharp denticle on the lateral margin and an anteriorly shifted basal pit, at which the basal margin is upturned.

5. Age assignment

Praemesosaturnalis heilongjiangensis was originally described in the Norian (Upper Triassic) in the Nadanhada Terrane, Northeast China (Yang and Mizutani, 1991). The species has also been reported from the Norian in other areas (e.g. Yao, 1982; Yoshida, 1986; Tekin, 2002). According to Sugiyama (1997), *Praemesosaturnalis*



Fig. 3 Locality map. The map is modified from the topographic map published by Geospatial Information Authority of Japan (https://maps.gsi.go.jp/).

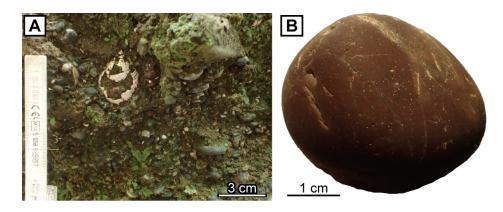


Fig. 4 Photographs of conglomerate and pebble. (A) Conglomerate of the Higashihigasa Formation of the Kazusa Group, Boso Peninsula. (B) Microfossil-bearing red chert pebble (sample 21112309a).

heilongjiangensis group mainly occurs in TR8A to TR8C, Norian, Upper Triassic.

The genus *Mockina* occurs in the middle to upper Norian (upper Alaunian to lower Sevatian) (Mazza *et al.*, 2012; Rigo *et al.*, 2018). The updated integrated biostratigraphy of radiolarians and conodonts by Yamashita *et al.* (2018) showed the co-occurrence of *Praemesosaturnalis heilongjiangensis* group and species of *Mockina* in the middle to upper Norian.

Based on the above-mentioned radiolarian and conodont occurrences, the sample is middle to late Norian in age.

6. Implication

In Southwest Japan, Triassic chert is a component rock of Jurassic accretionary complexes of the Tamba–Mino– Ashio and Chichibu belts (e.g. Nakae, 2000; Matsuoka *et al.*, 1998; Kojima *et al.*, 2016). Consequently, the chert pebble dealt with in this study must be derived from one of these Jurassic accretionary complexes. Meanwhile, the Jurassic accretionary complex is not exposed near the Boso Peninsula (Fig. 6). The nearest-exposed Jurassic accretionary complexes are those in the Yamizo Mountains (Ashio Belt), Ashio Mountains (Ashio Belt) and Kanto

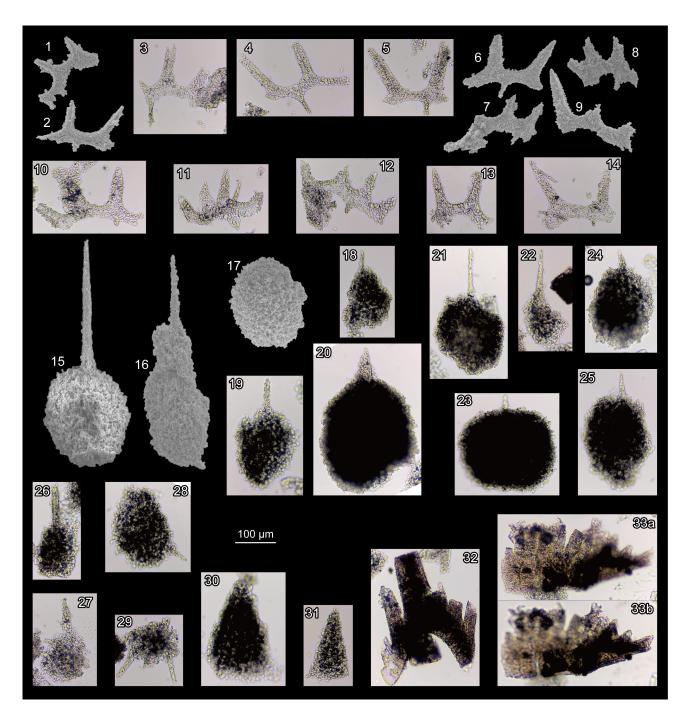
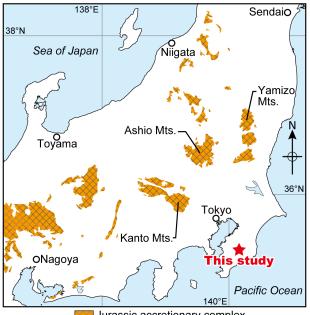


Fig. 5 Late Triassic radiolarians (1–31) and conodonts (32, 33) from the chert pebble (sample 21112309a). (1–5) *Praemesosaturnalis* sp. cf. *P. heilongjiangensis* Yang and Mizutani. (6–14) *Praemesosaturnalis*? sp. (15) *Paroertlispongus*? sp. (16–28) Spumellaria gen. et sp. indet. (29) Poulpidae? gen. et sp. indet. (30, 31) Nassellaria gen. et sp. indet. (32) Breviform digyrate conodont element (M element?). (33) *Mockina* sp.



Jurassic accretionary complex

Fig. 6 Distribution of Jurassic accretionary complexes in central Japan. The distribution is based on Geological Survey of Japan, AIST (2020).

Mountains (Chichibu Belt). The Jurassic accretionary complex in one of these mountains is the presumable origin of the chert pebble. Triassic chert pebbles have also been reported from the Neogene in the Boso Peninsula (Yamamoto *et al.*, 2012), so the chert pebble might be secondarily derived from such strata.

In the current knowledge, the origin of the pebble is not conclusive. Triassic chert is found throughout most of the Jurassic accretionary complex, whereas some lithologies (e.g. limestone, Permian chert, Upper Jurassic mudstone) are unevenly distributed in the complex. If clasts of these lithologies can be found, their origin can be determined in more detail. Further accumulation of the data of the Quaternary in and near the Boso Peninsula will clarify its origin more accurately.

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房総半島,上総層群の下部更新統東日笠層のチャート礫から産出した 後期三畳紀放散虫及びコノドント

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要 旨

房総半島の上総層群の下部更新統東日笠層に挟在する礫岩中のチャート礫から,放散虫及びコノドントが産出した. 放散虫 (*Praemesosaturnalis* sp. cf. *P. heilongjiangensis*) とコノドント (*Mockina* sp.)の同定に基づくと,このチャート礫は後 期三畳紀 (中期~後期ノーリアン期)の年代を示す.本チャート礫は当時後背地に分布していたジュラ紀付加体に由来す ると考えられる.